

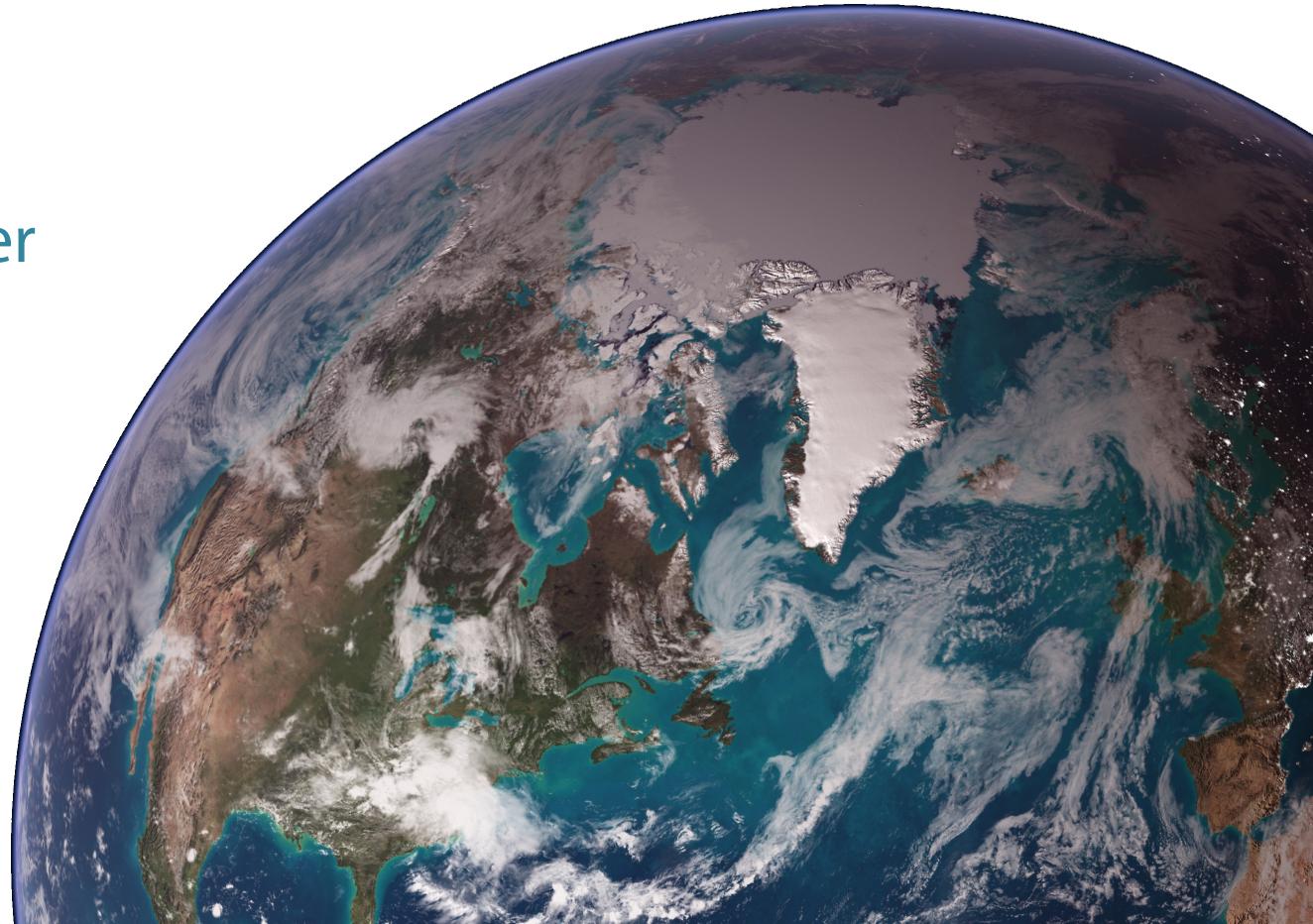


NOAA RESEARCH • ESRL • PHYSICAL SCIENCES DIVISION

# The Processes Underlying the Pacific Decadal Oscillation

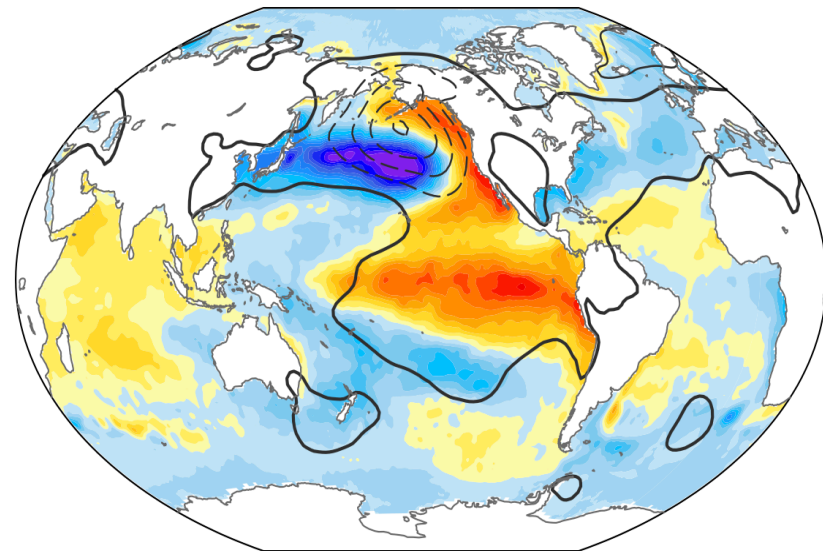
Michael Alexander

Science Review  
12-14 May 2015  
Boulder, Colorado

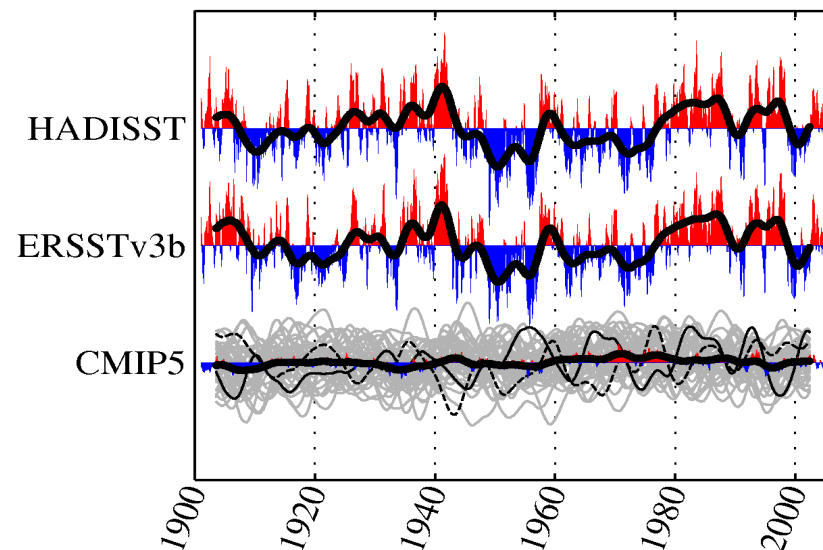


# Why study the PDO?

- Leading pattern of SST variability in the North Pacific ( $> 20^{\circ}\text{N}$ )
- Associated with climate, ecosystem and hydrologic fluctuations
- Develop a process understanding - key to prediction and applications

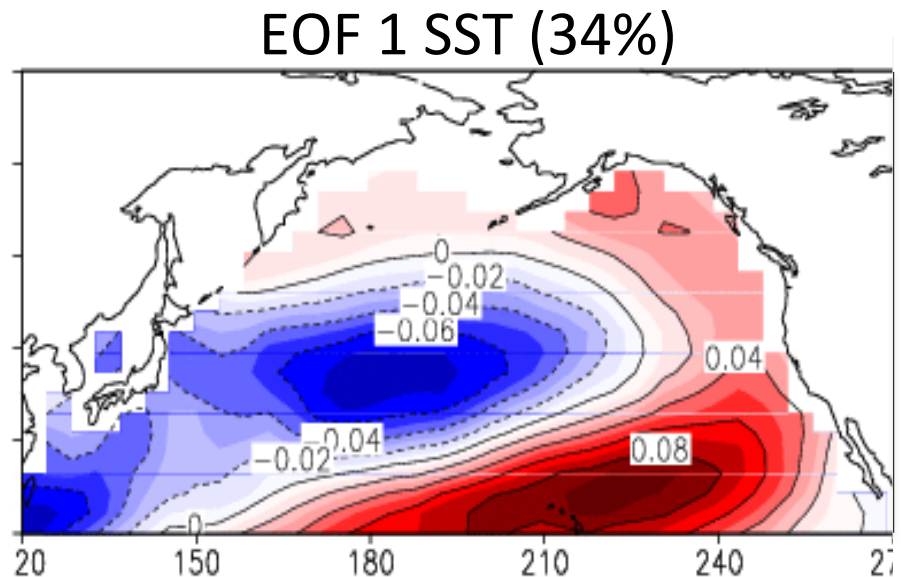
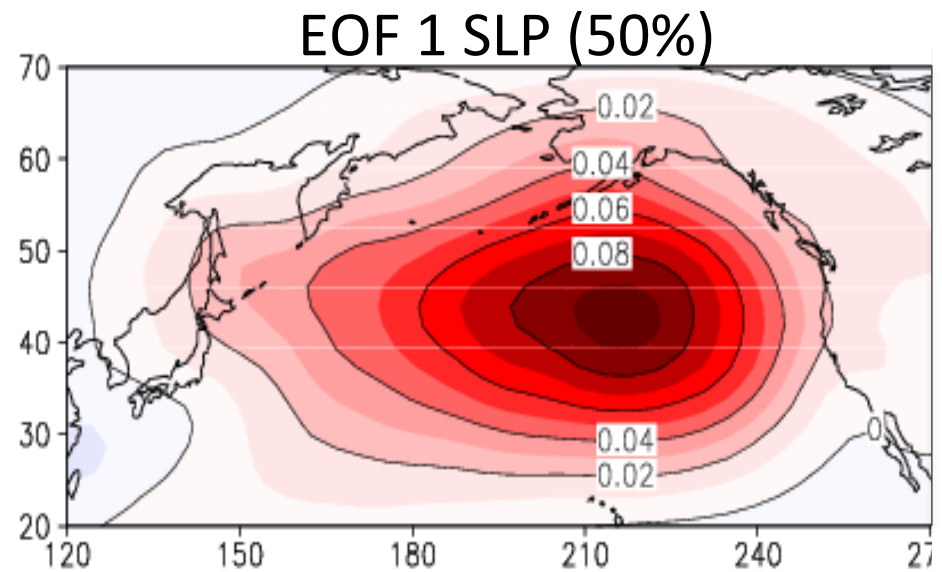


**c. PDO Time Series**



# Random Forcing Aleutian Low Variability

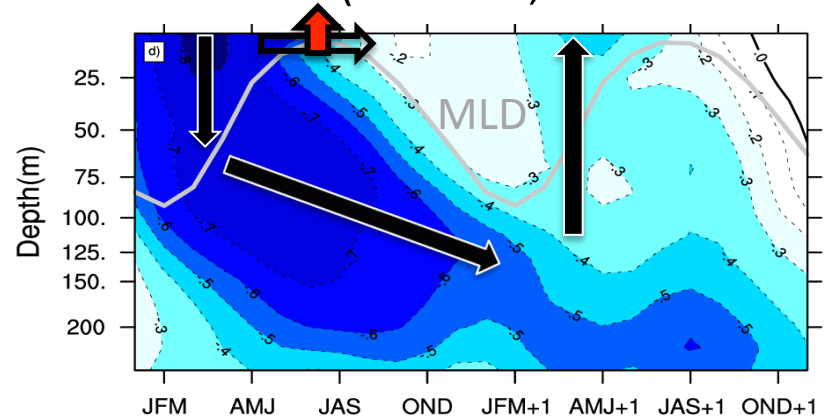
- Ocean is a simple slab no currents thus no ENSO or ocean gyres
- Leading pattern => changes in strength of the Aleutian Low
- Changes in surface fluxes forces ocean
- Ocean integrates flux forcing creates SST anomalies that resemble the PDO



# Midlatitude Ocean Processes

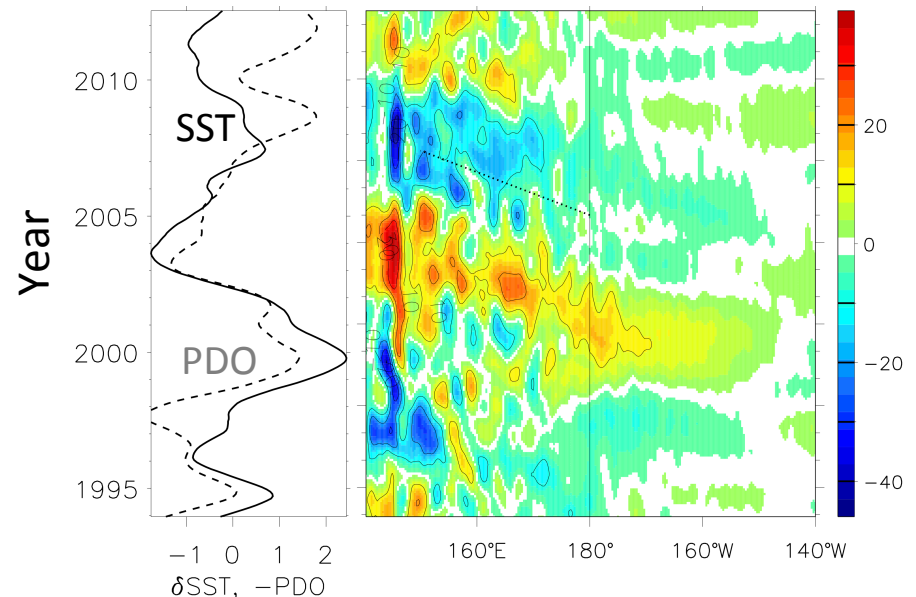
- Reemergence
  - Acts to “lengthen” ENSO & random atmospheric forcing
- Wind generated Ocean Rossby waves
  - Impacts SST near Japan, along the Kuroshio-Oyashio Extension (KOE) front

Central Pacific (35°-45°N, 170°W-150°W)



West

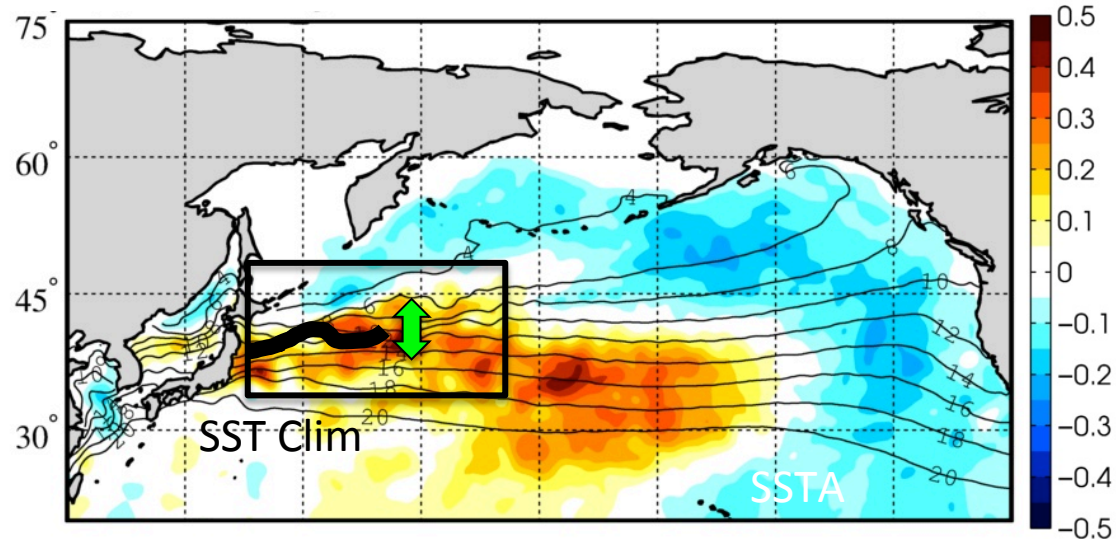
Pacific SST Sea Surface Height (33°-35°N)





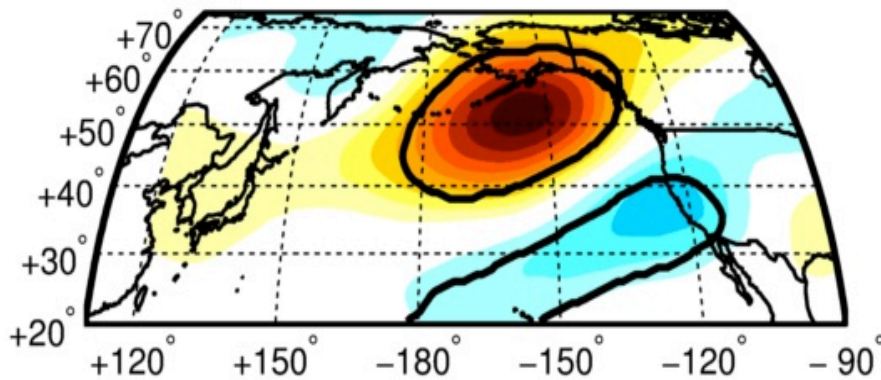
# Kuroshio-Oyashio frontal variability

- SST anomalies and the atmospheric response to the frontal anomalies in an atmospheric model

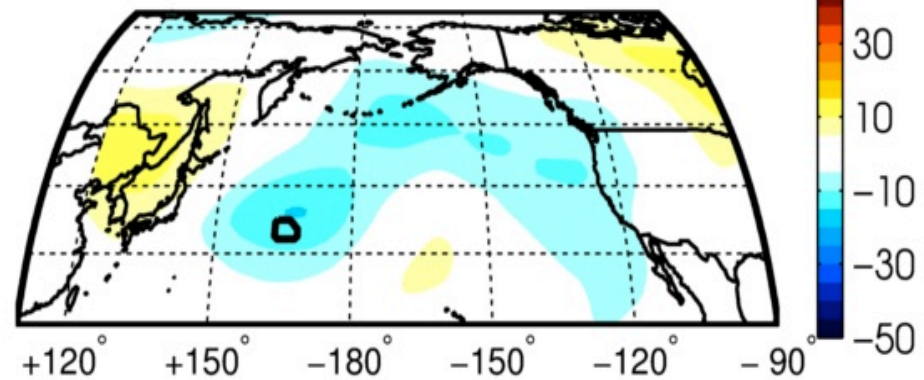


300 hPa height (m)

CAM5 1/4°

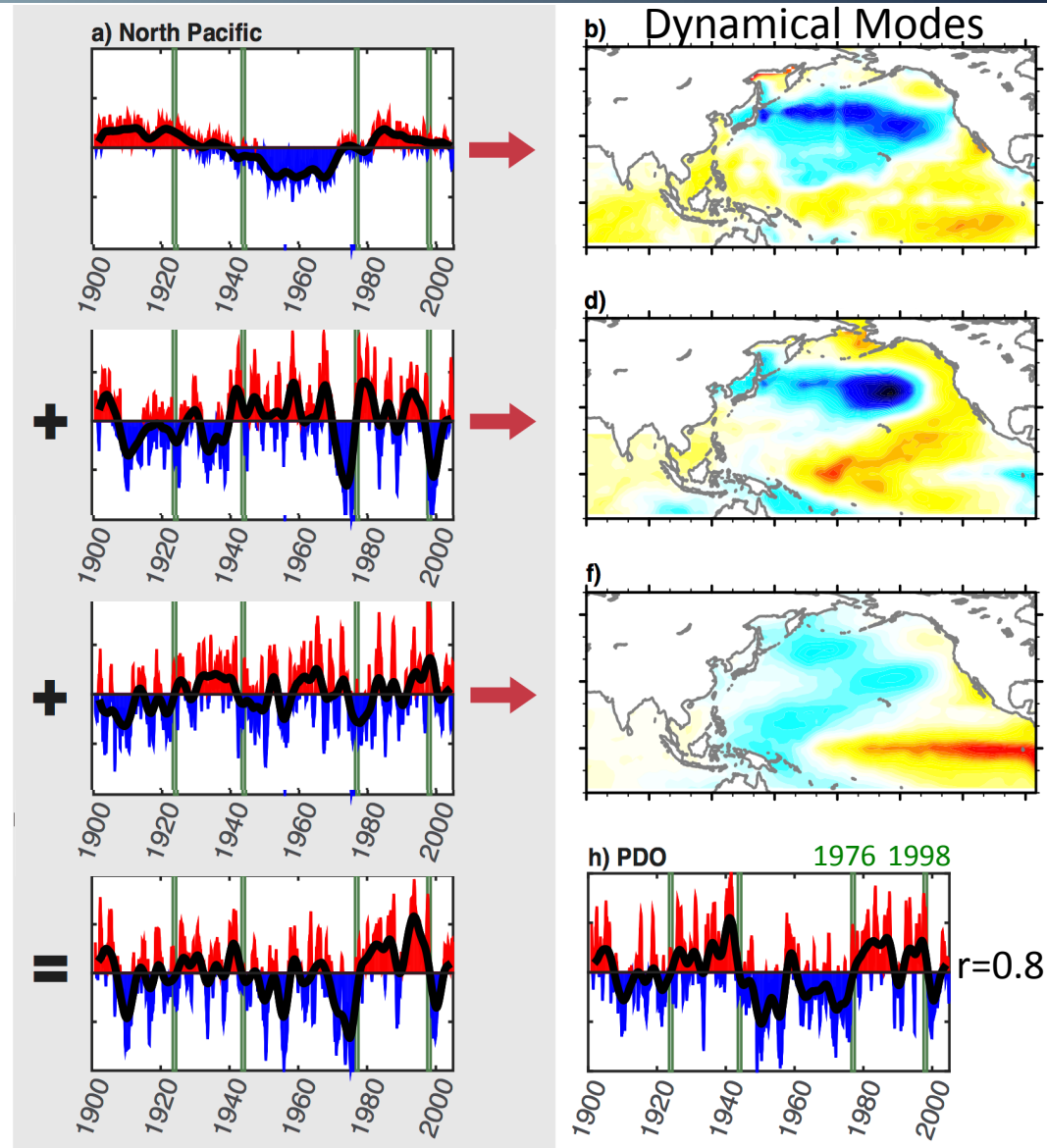


CAM5 1°



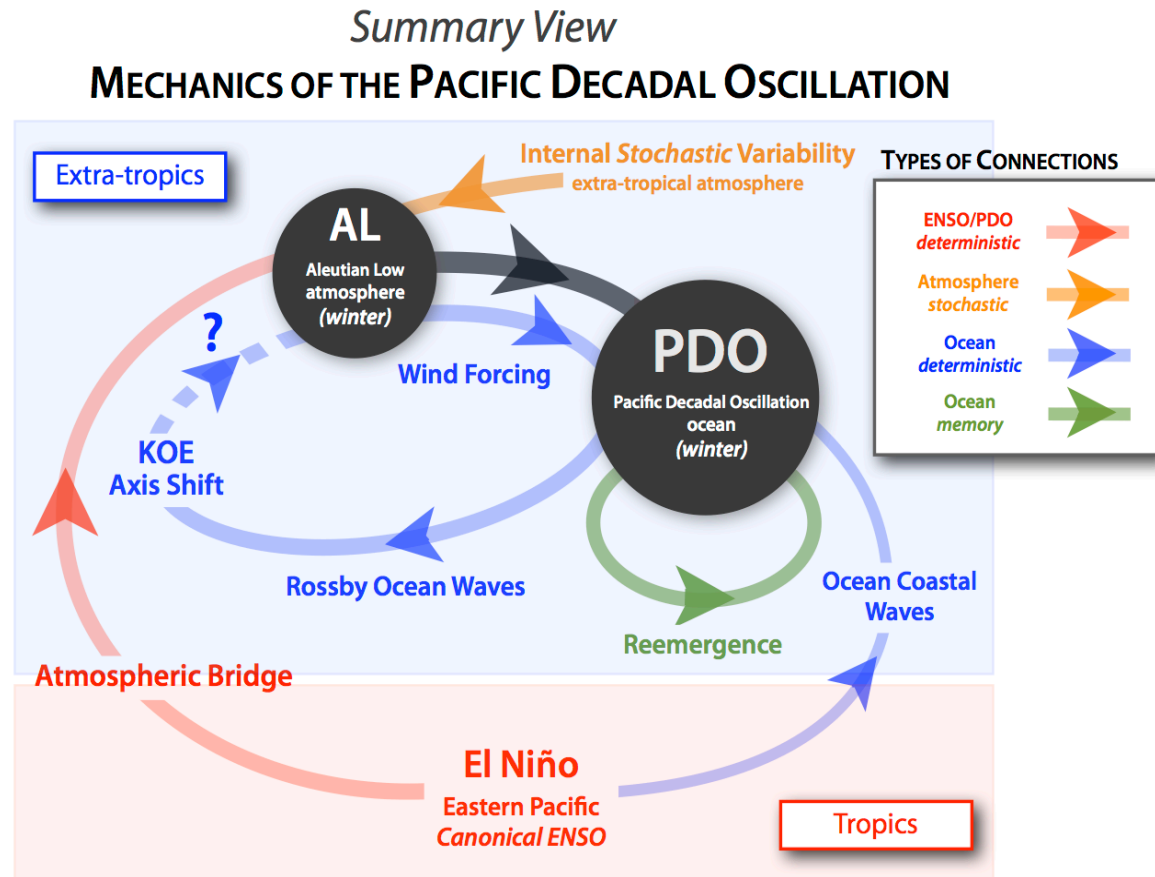
# Building the PDO

- Empirical Model (LIM)
- Leading Pacific dynamical modes
  - Not EOFs, not orthogonal
- Time series show projection of each mode onto the PDO



# Summary and Conclusions

- Processes:
- Atmospheric Bridge (ENSO)
- Random forcing
- Reemergence
- Ocean Rossby waves & ocean fronts
- Atmospheric response to KOE SST anomalies?



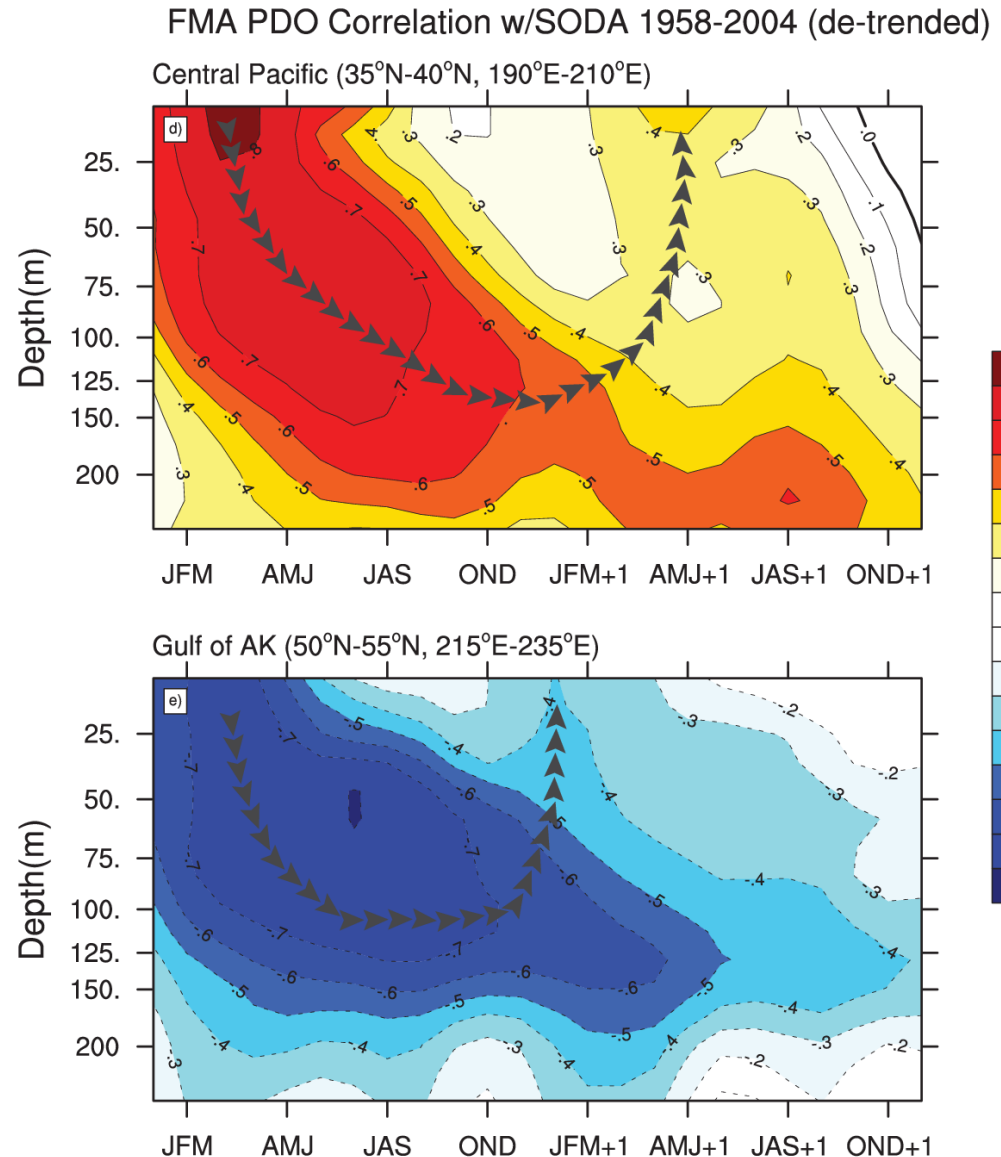
# Additional slides



“Re-emergence” : SST anomalies can recur in consecutive winters in the extratropics

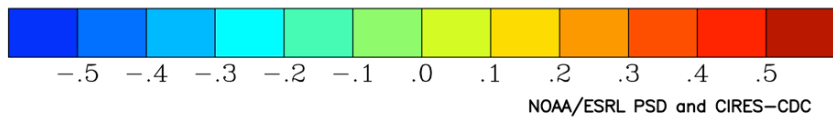
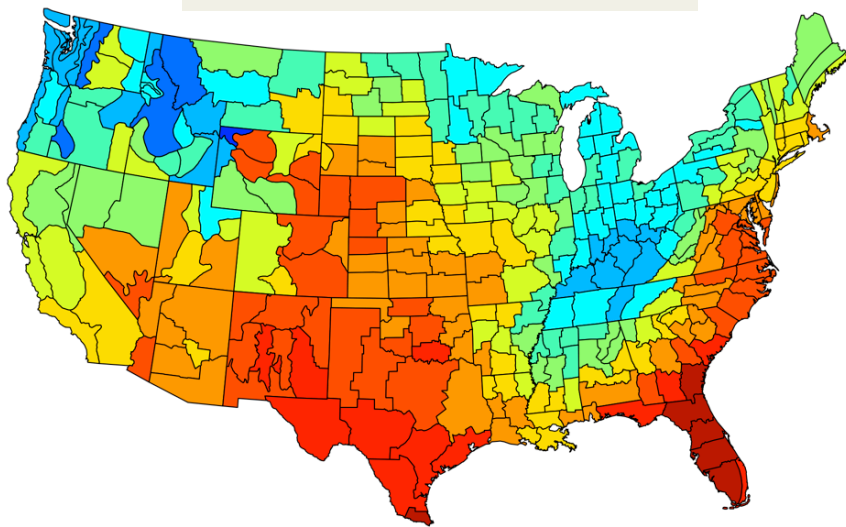
Acts to lengthen ENSO and Random Aleutian Low forcing

Depth vs. time cross-section of ocean temperature anomalies ( $^{\circ}\text{C}$ ) in two regions, correlated on the PDO (1958-2004)

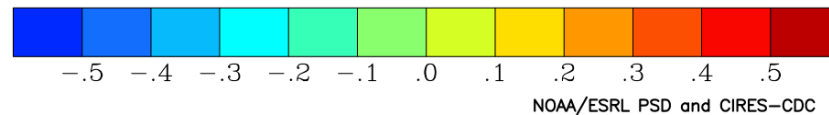
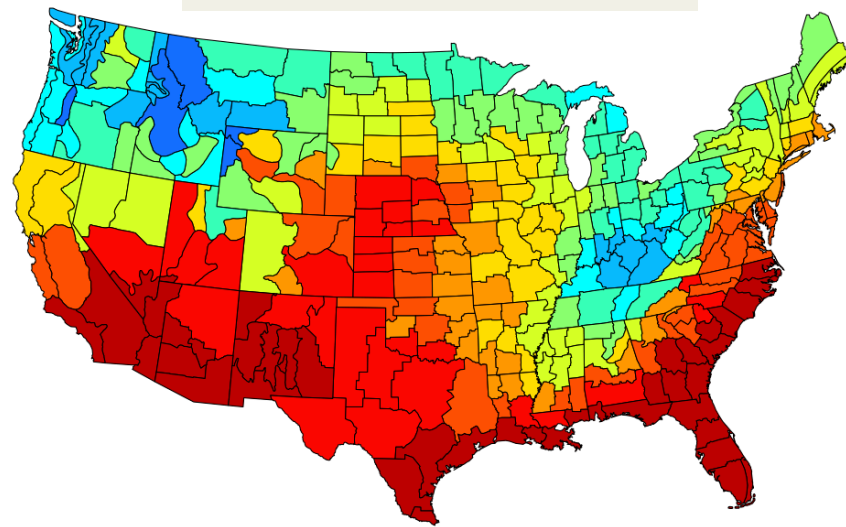


# PDO and ENSO "climate signals" are not independent

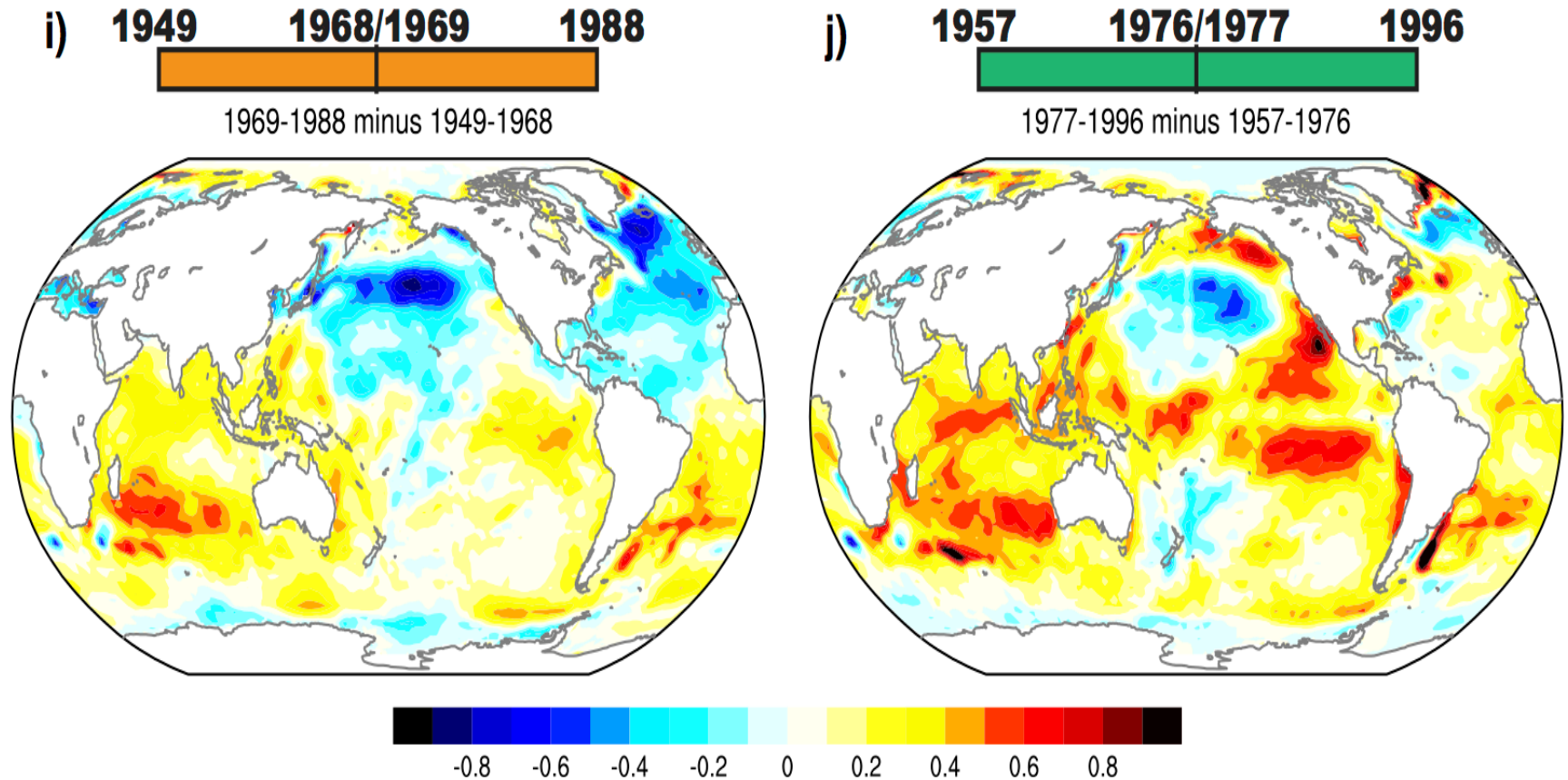
Nov-Mar precipitation correlated with PDO



Nov-Mar precipitation correlated with ENSO



# Epoch differences in SST



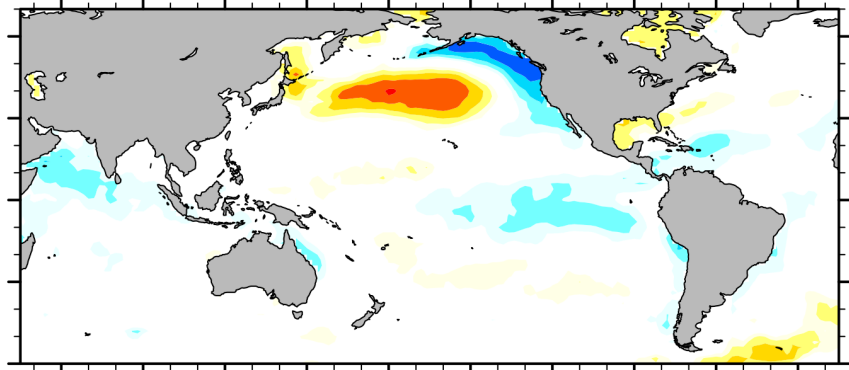
# Local and remote forcing of the PDO

Top: atmosphere (NPI) leads SST by three months

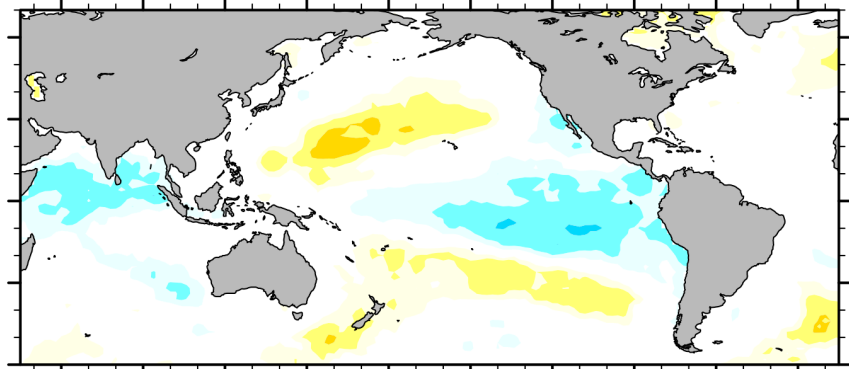
Middle: SST leads NPI by three months

Bottom: ENSO index leads SST by three months

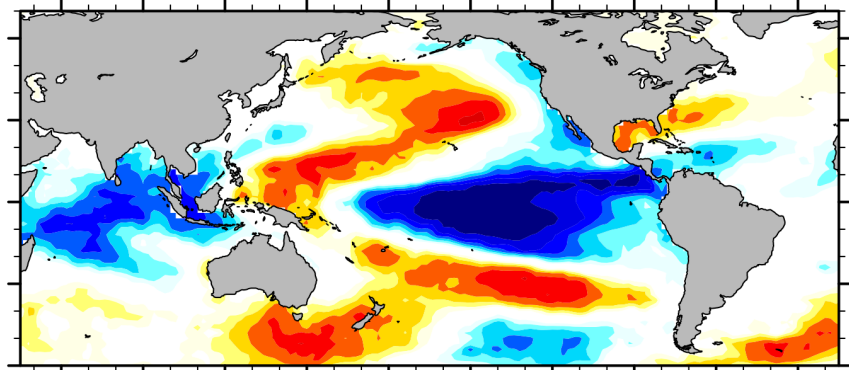
NDJ NPI leads SST by 3 months (correlated with FMA SST)



FMA NPI lags SST by 3 months (correlated with NDJ SST)



NDJ tropical Pacific -PC1 leads SST by 3 months (correlated with FMA SST; flipped sign)

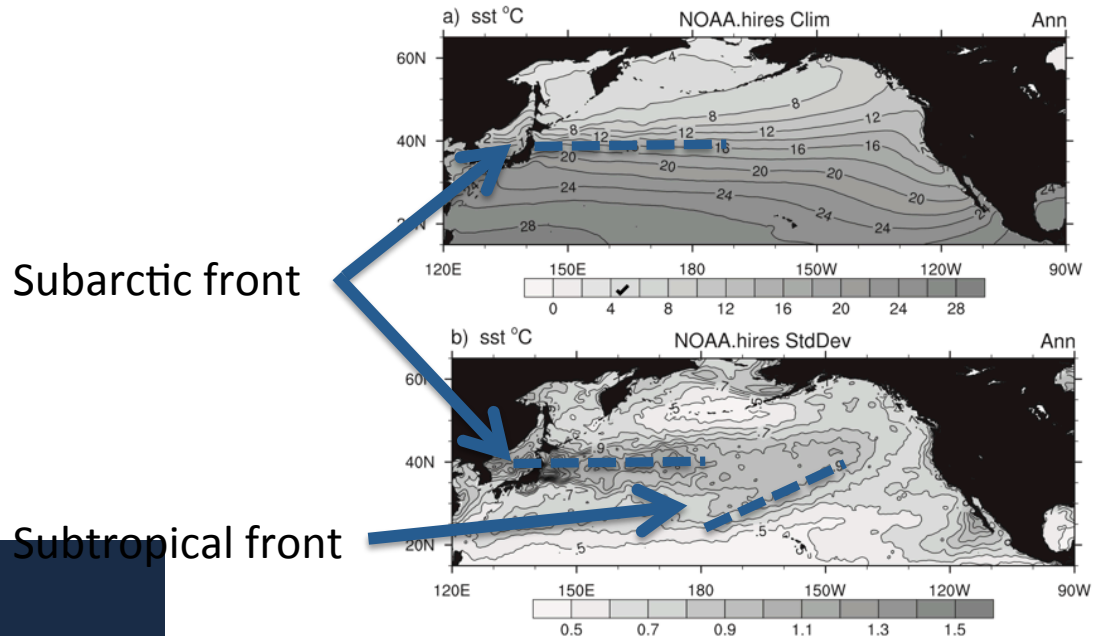
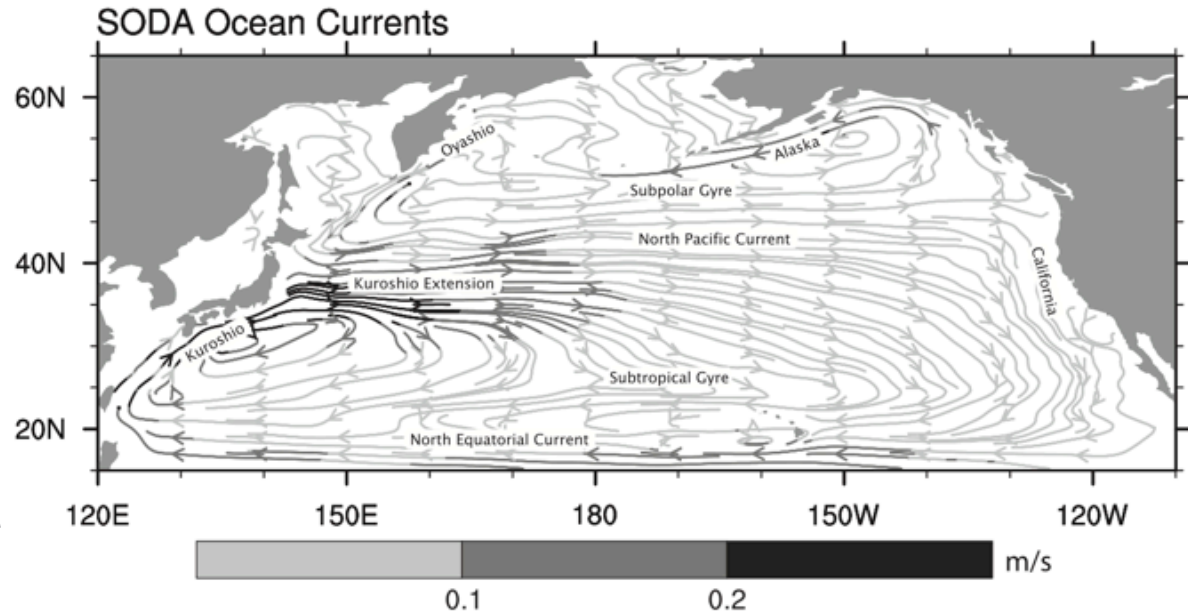




# Pacific Ocean currents and variability

Kuroshio-Oyashio Extension (KOE) system is a key component of the North Pacific ocean-atmosphere system

Shifts in the subarctic SST front are associated with longer time scales (westward propagating Rossby waves)



# Removing tropically-forced portion of the PDO yields “internal” North Pacific SST mode

Multivariate AR1 model (LIM):

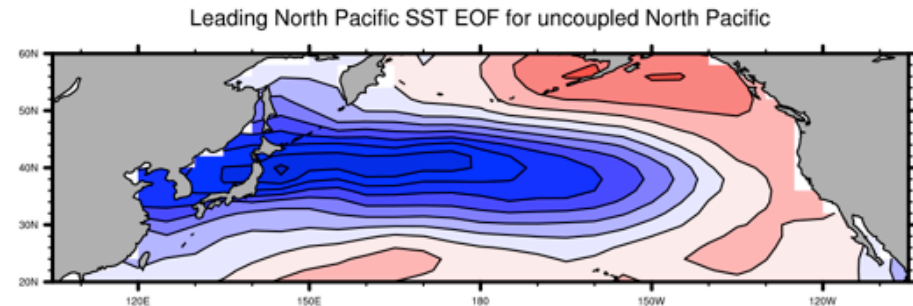
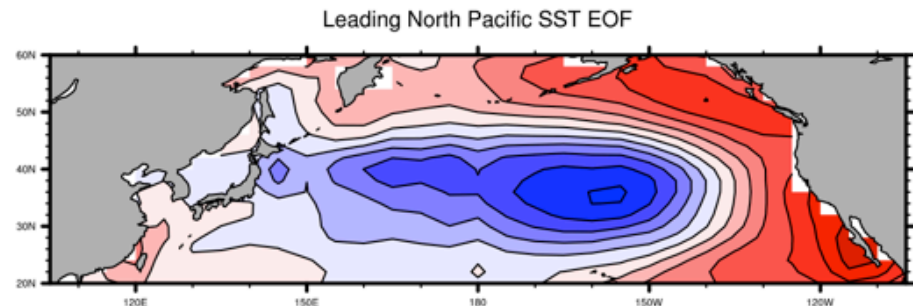
$$dx/dt = \mathbf{B}x + \mathbf{F}_s$$

Determined from observations, where  $\mathbf{x}$  represents seasonal mean anomalies (1958-2008) of

- **Tropical Pacific**  
[SST, thermocline depth]
- **North Pacific**  
[SST, mixed layer temp (30-100m)]

Top: Leading pattern of North Pacific variability (PDO)

Bottom: Leading pattern of “internal” North Pacific seasonal variability (after uncoupling Tropics and North Pacific dynamics within  $\mathbf{B}$ )



# ENSO-PDO representation in CMIP5

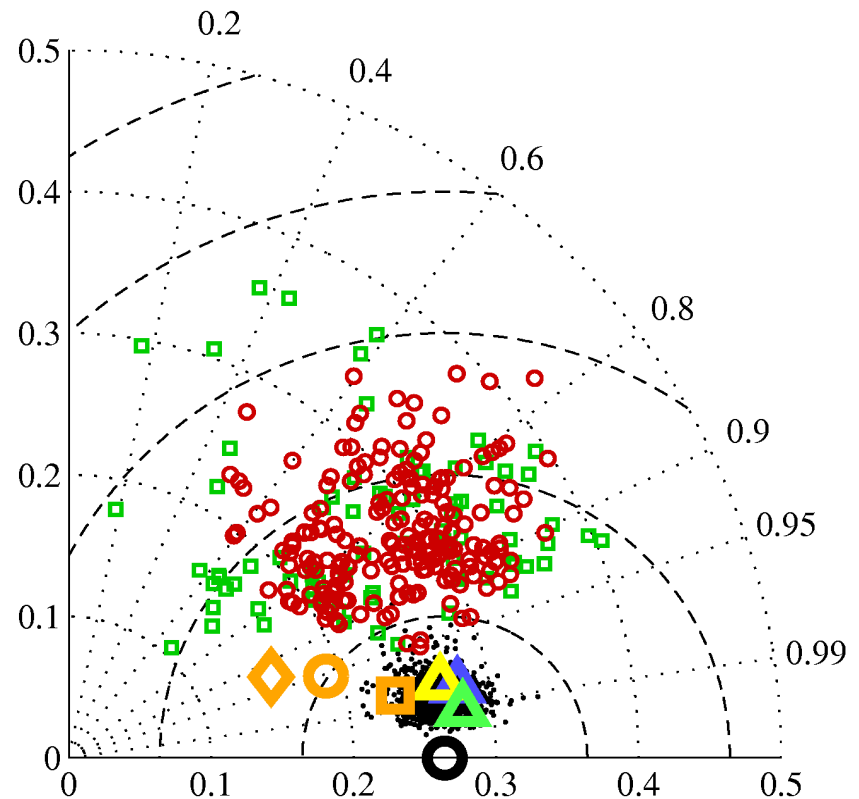
Taylor diagram compares PDO determined from HadISST, 1901-2004, to

- CMIP3 : green
- CMIP5: red
- Black dots: 50-yr Monte Carlo subsampling
- Triangles: other data sets

Key result:

- Models reproduce a PDO EOF but none reproduce PDO well

**c. PDO Taylor Diagram**



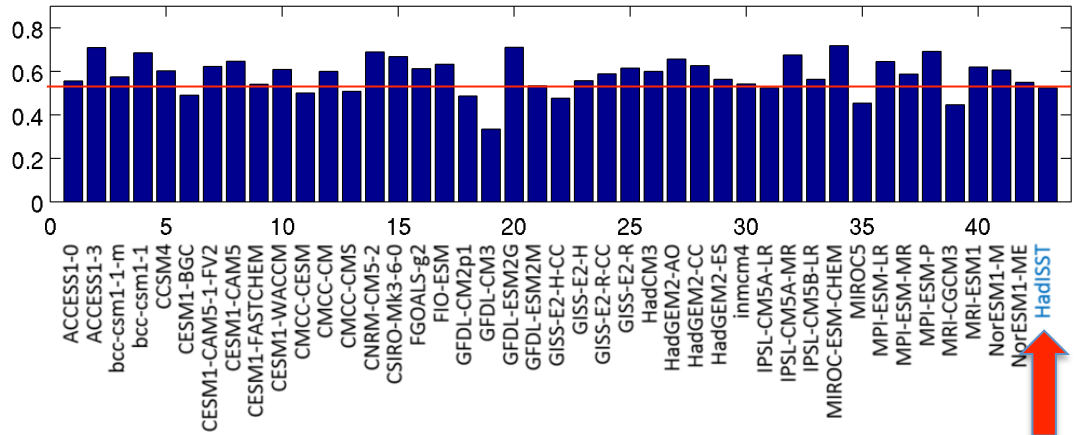
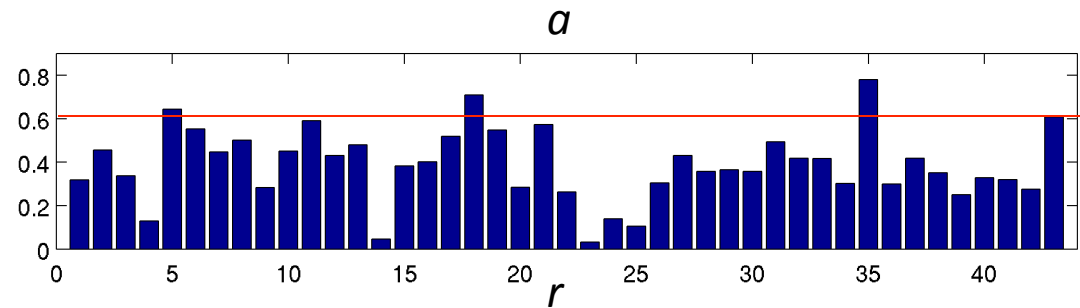
# ENSO-PDO representation in CMIP5

$$\text{PDO}(n) = r \text{PDO}(n-1) + a \text{PC1}_{\text{Tropics}}(n) + b \text{PC2}_{\text{Tropics}}(n) + e$$

Fitting (simpler) AR1 model to observations and CMIP5 models, 1901-2004

Key results:

- Most models reproduce PDO EOF
- Almost all models underestimate tropical forcing of PDO ( $a$ )
- Most models (slightly) overestimate  $r$

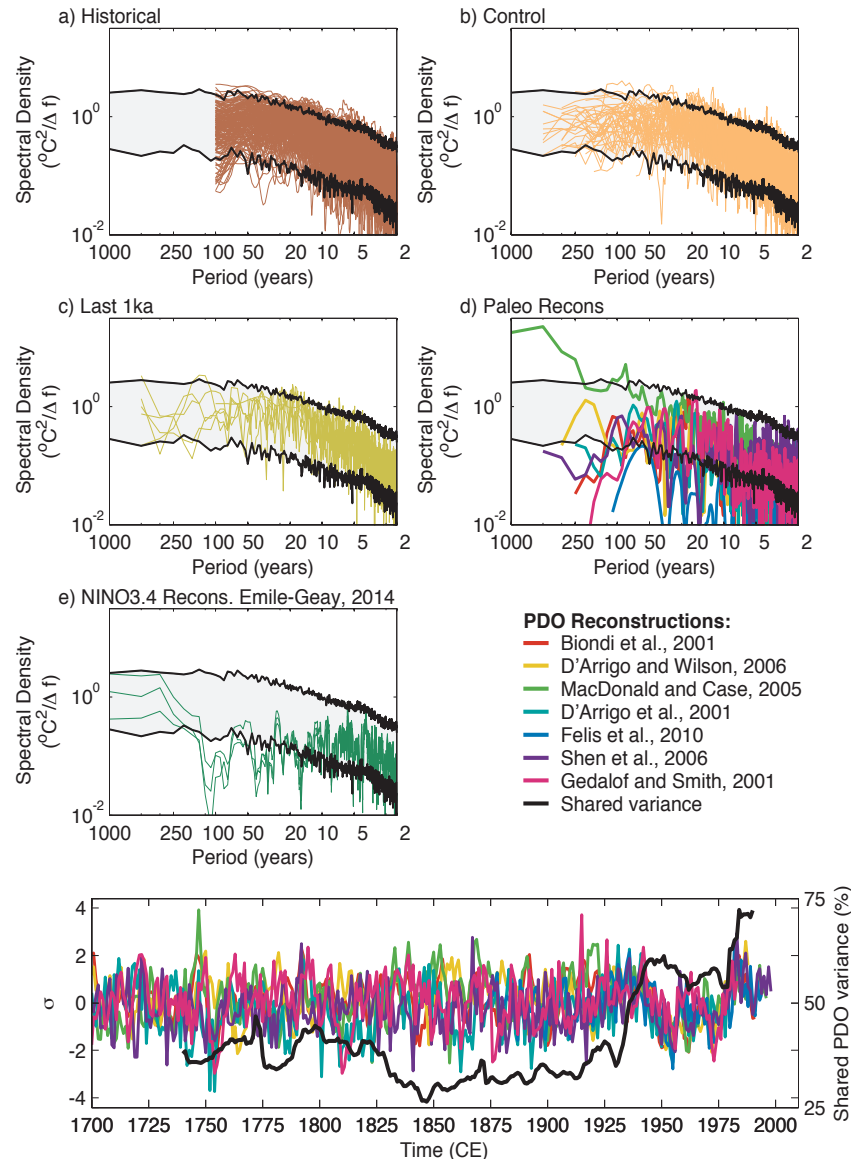




# PDO/ENSO spectra

Gray shading:  
1000 1000-yr LIM  
(multivariate AR1)  
realizations

CMIP5 spectra  
lies within  
confidence  
interval (a-c)



PDO paleo  
reconstructions