

AN EMPIRICAL ESTIMATE OF THE SOUTHERN OCEAN AIR-SEA CO₂ FLUX

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ABSTRACT

A discrepancy exists between current estimates of the Southern Ocean air-sea flux of CO₂. The most recent estimate using a combination of direct and climatologically-derived pCO₂ measurements [Takahashi *et al.*, 2002] (herein referred to as T02) suggests a Southern Ocean CO₂ sink that is nearly two times greater than that suggested from general circulation models, atmospheric inverse models [Gurney *et al.*, 2002] and oceanic inverse models [Gloor *et al.*, 2003]. Here we employ an independent method to estimate the Southern ocean air-sea flux of CO₂. Our method exploits all available surface measurements for Dissolved Inorganic Carbon (DIC) and total alkalinity (ALK) from 1986 to 1996. We show that surface age-normalized DIC can be predicted to within ~8 μmol/kg and ~10 μmol/kg for ALK using standard hydrographic properties, independent of season. The predictive equations are used in conjunction with World Ocean Atlas (2001) climatologies to estimate an annual cycle of DIC and ALK, while the pCO₂ distribution is calculated using standard carbonate chemistry. For consistency we use the same gas transfer relationship and wind product from Takahashi *et al.*, [2002] however, we include the effects of sea-ice. We estimate a Southern Ocean CO₂ sink (>40°S) of -0.19±0.26 Pg C for 1995. Our estimates are smaller than those estimated by Takahashi *et al.*, [2002], but consistent with atmospheric / oceanic inverse methods, general circulation models and provides further evidence that the Southern Ocean CO₂ sink in relation to its oceanic surface area, is moderate on a global scale.

DATA AND METHODOLOGY

Our methodology to estimate the air-sea CO₂ flux in the Southern Ocean follows a stepwise procedure. We firstly develop a predictive empirical relationship for DIC and ALK that uses all available carbon measurements over the past two decades where DIC has been normalized to a common year (1995) using CFC measurements. A step-wise linear least squares regression for surface DIC and ALK was conducted using various independent parameters as predictors. The regression equation for DIC is represented by:

$$DIC_{obs} = \alpha_0 + \sum_{i=1}^n \alpha_i P_i + \varepsilon, \text{ where } \alpha_i \text{ are the partial regression coefficients for } n \text{ number of independent}$$

parameters (P_i) and α₀ is the y-intercept. The independent parameters were varied to investigate the influence of each on the DIC prediction. Each regression was applied to the upper 20m of data south of 40°S. The final resulting fit for DIC uses 5 parameters (T, Sal, O₂, NO₃ and SiO₄) with a standard error of ~8.4 μmol/kg and an adjusted R² of 0.97. We then exploit the predictive relationship by using all available hydrographic properties from the World Ocean Atlas 2001 in order to estimate an annual cycle of DIC and ALK for the year 1995. Using optimal dissociation constants, the annual cycle of pCO₂ for 1995 is then calculated. The final step was to use global wind products and gas exchange velocities to estimate a net uptake of CO₂ in the Southern Ocean and to quantify the effects of sea-ice. A careful uncertainty analysis using Monte-Carlo trials precludes a final estimate of CO₂ uptake in the Southern Ocean and a comparison to other independent results.

RESULTS

In general, our pCO₂ estimates are higher than those estimated from T02 (Fig. 1), with the only exception being within the sub-Antarctic Zone in winter. Spring, autumn and winter pCO₂ estimates within the SAZ are in good agreement within T02 when considering the uncertainties. Regions showing significant differences are north of 65°S in summer and south of 55°S in the winter. Overall our flux estimate (~ -0.2 Pg C /yr) implies a weaker Southern Ocean CO₂ sink than from T02 (~ -0.35 Pg C /yr). New pCO₂ measurements within the Indian sector of the Southern Ocean have also challenged the notion of a large Southern Ocean CO₂ sink [Metzl *et al.*, 2004]. Metzl *et al.*, [2004] take seasonal pCO₂ measurements in the year 2000 and extrapolate using the 1-D biogeochemical model of Louanchi and Hoppema [2000] and find a considerably weaker CO₂ sink south of the Polar Front (~0.1 Pg C / yr). They suggest the weaker CO₂ sink in the Southern Ocean is due to the inclusion of direct winter-time pCO₂

measurements. Although this result is different to T02, it is consistent with our predictions and other independent estimates (Table 1).

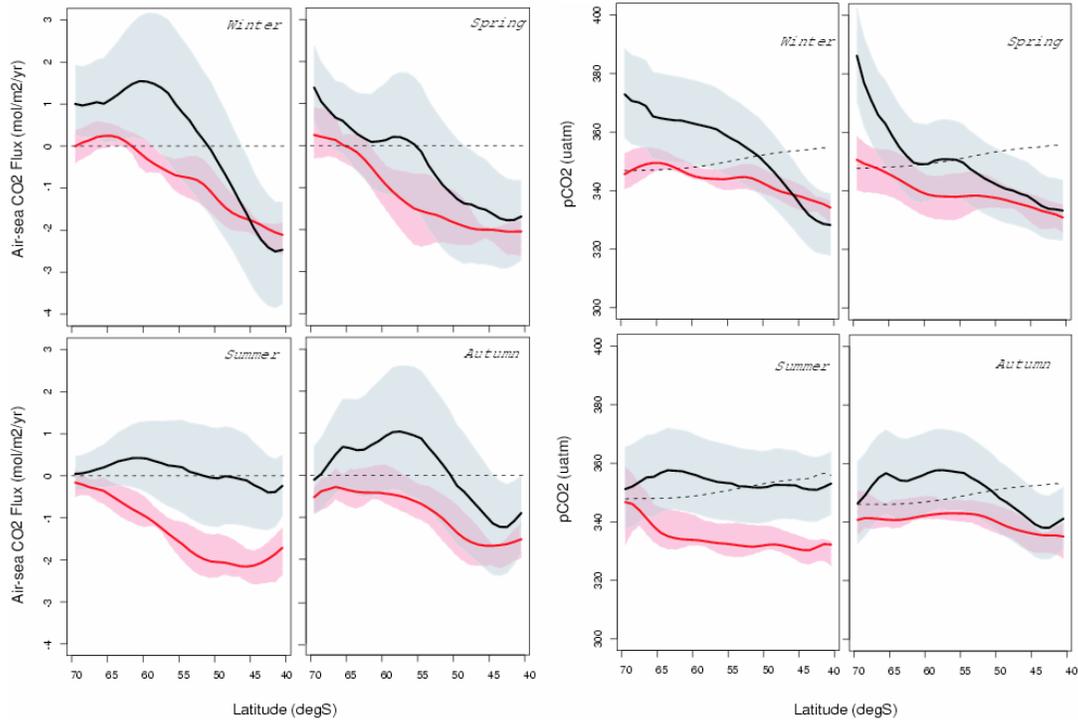


Fig. 1: Zonally average pCO₂ (µatm) and air-sea flux distributions for each season in the Southern Ocean. The red line and associated shading shows the T02 estimates. The solid black line and grey shading shows our estimates and associated uncertainty. The dashed line represents the atmospheric pCO₂ concentrations in 1995.

Table 1: Recent estimates of the air-sea CO₂ flux in the Southern Ocean

Reference	Method and Latitude Band	Estimate (Pg C /yr)
(Gloor et al., 2003)	Oceanic Inversion (>58°S)	-0.1
(Gurney et al., 2002)	TRANSCOM-2 Atmospheric Inversions (>45°S)	-0.4±0.3
(Takahashi et al., 2002)	Oceanic pCO ₂ climatology (>50°S) – NCEP 0.995 sigma winds	-0.5 to -0.7
Takahashi (2004)	Oceanic pCO ₂ climatology (>50°S) – NCEP 10m winds	-0.35 to -0.45
(Louanchi and Hoppema, 2000)	Semi-prognostic Model with some observational constraints (1984-1994)	-0.5
(Metzl et al., 2004)	Summer/winter pCO ₂ measurements (>50°S)	-0.1
This Study	Oceanic DIC/ALK climatology (>40°S)	-0.2±0.3

References

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